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# ANALYSIS ON THE NEW CHARACTERISTICS OF WIND WAVES AT SEA SURFACE

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#### ABSTRACT

In order to comprehend the characteristics of sea waves, starting from the features of fluid dynamics, the relationship between wave height and sea surface wind speed was found, the wind speeds over sea surface were calculated with SMB formula, so got the corresponding relationship between the peak frequency of wave height and the wind speed. Sea wave spectrum was introduced, the simulation results suggested when the wave number is smaller, changing of wave velocity has greater impact on sea wave spectrums. By using the resonance condition between radar working frequency and that of sea wave, sea surface wave number model was set up. Simulation results revealed that although the cutoff wave number changes with wind speeds and radar frequencies, but in the range of normal working frequencies of HF radar and general sea wind speeds, the cutoff wave number basically stabilizes as a constant quantity. Research conclusions can provide theoretical basis for quantitative research of the sea waves characteristics; it also is an important basis for the simulation of radar sea echoes.

Keywords: Wind Waves, Sea Wave Spectrum, Wave Height, Resonance Condition, Cutoff Wave Number

#### 1. INTRODUCTION

Wind acting on the sea surface to produce wave, and the wind's continuing role would form waves of the sea. According to Fourier theory, the sea surface can be regarded as a superposition of waves of different amplitudes, phases, frequencies and directions of arrival. When the effect of wind on sea continues for some time, the energy which sea waves derives from the wind could compensate for the energy consumption by air resistance, the sea surface reaches a state of dynamic equilibrium - steady state, and now the waves heights only relevant with the wind speed; but not all of the surfaces are unconditionally to achieve stable state. There are certain limitations to quantitative study of sea wave characteristics [1], the research needs to be further strengthened.

The time-varying characteristics of sea surface waves are very complex; it can only be described by using statistical methods. According to Fourier theory, any surfaces including those completely random surfaces can be represented as a superposition of different amplitudes, phases, frequencies and wave arrival directions; these waves should meet the dynamics theory.

The deep-water gravity wave is the main object of study. Typically, if the depth of sea wave is greater than half the wind wavelength, the condition of deep water wave is considered satisfied.

To a steady sea surface, the sea wave height h can be expressed as [2]:

$$h(x, y, t) = \sum_{n=1}^{\infty} a_n (k_x x \cos \theta_n$$
(1)  
+ $k_n y \sin \theta_n - \omega_n - \psi_n$ )

Of which,  $a_n$  is the single-wave amplitude,

 $\omega_n$ ,  $k_n$ ,  $\theta_n$  and  $\psi_n$  the angular frequency, wave number, propagation direction angle and initial phase respectively, meanwhile  $0 \le \theta_n \le 2\pi$ , probability distribution function  $\psi_n$  being uniform distribution, i.e.

$$p(\psi) = \frac{1}{2\pi}.$$
 (2)

The wave height distributions to steady sea surface belong to Gaussian distribution:

$$p(h) = \frac{1}{\sqrt{2\pi}\sigma_n} \exp(-\frac{h^2}{2\sigma_n^2}), \qquad (3)$$

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Where  $\sigma_n^2$  is the variance of sea wave height *h*.

Although various wavelengths are included in the sea wave height of formula (1), but the most important are the gravity wave and capillary wave, the gravity one is our key research object.

#### 2. CHARACTERISTICS OF SEA WAVES HYDRODYNAMICS

For deep-water gravity waves in steady state, there are the following relationships, which sea waves must comply with:

When sea surface getting to its steady state, the height of sea wave reaches its maximum [3]

$$H_{\rm max} = \frac{0.26}{g} v_{10}^2 \,. \tag{4}$$

Where  $v_{10}$  is the wind speed 10m above the sea surface,  $g(m/s^2)$  is the gravitational acceleration.

Characteristic velocity of sea wave:

$$v = \sqrt{\frac{g}{k}} = \sqrt{\frac{g\lambda_s}{2\pi}} = \frac{g}{\omega}$$
(5)

Gravity wave frequency:

$$f = \frac{v}{\lambda_s} = \sqrt{\frac{g}{2\pi\lambda_s}} \tag{6}$$

Sea wave and its wavelength:

$$\lambda_s = \frac{g}{2\pi f^2} \tag{7}$$

Sea wave and its wave number:

$$k_s = \frac{4\pi^2 f^2}{q}.$$
 (8)

#### 3. CALCULATION OF SEA WAVE PARAMETERS - SEA WAVE SPECTRA

Sea wave spectra is the wave energy spectra, it gives directly the relationship between sea wave energy, wave frequency and wave direction. Because it can be easily obtained by observation, so it is the most fundamental method to describe sea waves.

Such as the adopted sea spectrum in actual use the Pierson-Moskowitz spectrum [4, 5, 6] (PM), below is its undirected energy spectrum,

$$s(k) = \left(\frac{\alpha}{2}\right) k^{-4} \exp(-b(k_e / k)^2)$$

$$k_e = \frac{g}{v^2}, b = 0.74, v(h \ge 10m).$$
(9)

Where,  $\alpha$  is Phillip constant quantity.

Figure 1 shows the undirected energy spectrum corresponding to velocity v of 4, 8, 12, 16 and 30m/s separately. Both horizontal and vertical coordinate axes of the figure choose the logarithm magnitudes.



Figure 1: The Relationship Between Sea Wave Energy Spectrum Of PM Model S (K) And The Sea Wave Number K

Simulation results show that for the PM wave: 1) the energy spectrum specific to different speed follows the same rule - It increases first and then decreases as wave number k keeping on increasing. 2) When with a higher speed, the wave energy spectrum has a larger maximum and mean value. 3) As the logarithm of wave number greater than 0.15, changing in wind speed has no differences to energy spectrum.

#### 4. CALCULATION OF THE WIND SPEED ABOVE SEA SURFACE

For different heights above sea surface, the air density would vary with the heights, the wind pressure on the sea is also different from each height; the damping coefficient which atmosphere acting on wind is different too [7], so the wind speeds are not same even being over the same sea-surface. Figure 2 was the vertical distribution of longitudinal wind speeds at a vertical cross section BB' [8].

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The relation between wind speed and the height H above sea is as [9]:

$$V_{H} = V_{10} \{ 1 + [(C_{10})^{1/2} / k] \ln(H_{E} / 10) \}$$
(10)

Where  $V_H$  and  $V_{10}$  are the wind velocities at H(m) and 10m above sea surface apart, k being Karman constant,  $C_{10}$  damping coefficient 10m above sea, and it having the relationship with  $V_{10}$  as the following [10]:

$$C_{10} = (0.80 + 0.11V_{10})10^{-3}$$
(11)

From the above two formulas, as long as  $V_{10}$  being established, the wind speed at any height above the sea can be determined.

The sea wave has two feature parameters; one is the valid wave height, the other the peak frequency of wave height spectrum. Both of these characteristic quantities are available through inversion of the second order radar cross section equation of sea echo Doppler spectrum.  $V_{10}$  can be computed by using SMB relational expression [11]:

$$\frac{gH_E}{V_{10}^2} = 0.26 \tanh[(f_M V_{10})^{-3/2} \frac{(3.5g)^{3/2}}{10^2}], \quad 12)$$

Where g is the gravitational acceleration, the numerical solution to  $V_{10}$  in transcendental equation (12), which is not easy to find, iteration method is used here; when solving it, first assuming the initial value of  $V_{10}$ , here setting 7, and substantially only iterating a few times, the result is stabilized. In addition, as sea waves needing to remain not broken, the effective height  $H_E$  and the

wavelength  $\lambda$  are demanded to meet  $\frac{H_E}{\lambda} \leq 1/7$ 

[12].

The second radar cross section equation of Doppler spectrum of sea surface echo is used to invert the wave height parameter [6, 13, and 14], formula (12) is employed to calculate wind speed  $V_{10}$ . Computed results as shown in table 1, from it can be got the conclusions: the peak frequency of wave height spectrum  $f_m$  changes only depending on wind speed quantity, and no relation to the wind direction. The greater, the wind speed, the smaller  $f_m$  is.

Table 1: Peak Frequency  $f_M$  , Effective Wave Height,

 $H_{\scriptscriptstyle E}\,$  and the Calculated Wind Speed  $V_{\scriptscriptstyle H}\,$ 

$f_{M}$	0.08	0.146	0.275
<sup>o</sup> M /Hz	0.08	0.146	0.275
Н	5.50	2.32	0.28
$II_E /m$	2.67	1.54	0.34
<b>V</b> 1	15.1185	10.5892	3.2786
$V_{H/m} \cdot s^{-1}$	10.0707	7.9275	3.6459

#### 5. SEA WAVES CUTOFF WAVE NUMBER OF RADAR

The cutoff wave number of sea wave is specific to narrow-beam radar. Radar Frequency and that of sea waves must meet the resonance condition. From formulas (5) and (6), a corresponding wave number  $k_c$  can be found, the wave number  $k_c$  is

normalized by using 
$$K_C = \frac{k_c}{2k_0}$$
, so got:

$$K_{C} = \frac{g}{v^{2}} \cdot \frac{1}{2k_{0}} = \frac{g}{v^{2}} \cdot \frac{c}{4\pi f_{0}}.$$
 (13)

The normalized cutoff wave number alters with the velocity v; it is also associated with the radar frequency  $f_0$ . This relationship is described in figure 3, figure 3\_a only reflects the pure mathematical relationship, figure 3\_b reflects a specific relationship which being in usual sea wind, normal working HF radar frequency.

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A: This Figure Only Reflects The Mathematical Relationship





By figure 3\_b, when in HF radar operating frequencies and general sea wind speeds,  $K_c$  fluctuates very little, so is applied as a constant. Of course, in specific case, sea surface wind speed is also relevant to the sea state; figure 3\_b does not reflect this meaning.

#### 6. CONCLUSION

Using SMB formula to calculate wind speed, the peak frequency of wave height spectrum changes only depending on wind speed quantity, the greater, the wind speed, the smaller the peak frequency. Simulation results showed: when the logarithm of wave number k is much smaller than 0.15, the sea wave energy spectrum increases significantly with the increasing of sea wind speed, but when the logarithm of greater than 0.15, wind speeds cause no diversities..

The quantitative research on the sea wave characteristics is mainly organized on the basis of measurement. With the measured HF radar sea echo data and the gradual accumulation of data, the establishment of a more practical model will be important research tendency in the future.

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#### **REFRENCES:**

- M. Feng , W. Y. Sha, Y. Li, Y. B. Hu. Development of Study on Waves Off shore. Journal of PLA University of Science and Technology, Vol. 5 No. 6, PP.70-76, 2004.
- [2] Y. R. Zhu. Wave Mechanics for Ocean Engineering[M]. Tianjing: Tianjin University Press, 1991.
- [3] D. Y. Huang. Character of HF Radar Sea Echo Spectra and Factor of the Effects on the Quality of the Spectra. Chinese Journal of Radio Science, Vol. 11 No. 2, PP. 94-101, 1996.
- [4] W. J. Pierson, W. Marks. "A proposed spectral form for fully developed wind seas based on the similarity theory of S. A. Kitaigorodskii. J Geophys. Res.", Vol. 69 No. 24, PP. 5181-5190, 1964.
- [5] L. Moskowitz. Estimatates of the power spectrums for fully developed seas for wind speed of 20 to 40 Knots, J.Geophys. Res., Vol. 69 No. 24, PP. 5161-5179, 1964.
- [6] Q. Z. Lei, J. X. Wang. Numerical Simulation and Analysis for the Sea Echo Spectrum of HF Radar. Journal of Xi'an Technological University, Vol. 28 No. 6, PP. 557-559, 2008.
- [7] Z. I. Rizman, K. Jusoff, S. S. Rais, H. H. H. Bakar, G. K. S. Nair, Y. K. Ho. Microwave signal propagation on oil palm trees: measurements and analysis. International Journal on Smart Sensing and Intelligent Systems Vol. 4, No. 3, PP. 388-401, 2011.
- [8] Y. G. WANG, Z. M. WU, Z. Q. CHANG. Numerical simulation of three dimensional structure of sea-land breeze at west coast of Liao Dong Bay. MARINE FORECASTS, Vol. 26 No. 3, PP. 12-21, 2009.

30<sup>th</sup> November 2012. Vol. 45 No.2

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[9] W. J. Pierson. The interpret	ation of wave	
spectra in terms of the wind pr	ofile instead of	
the wind measured at a c	onstant height.	
J.Geophys Res, Vol. 69 No. 24,	PP.5191~5203,	
1964.		

- [10] Q. Z. Lei, J. X. Wang. Extraction of wind field from radar sea-echo spectra. Journal of Zhengzhou University of Light Industry(Natural Science), Vol. 24 No. 1, PP. 121-124, 2009.
- [11] P. E. Dexter and Theodoridis. Surface wind speed extraction from HF sky wave radar Doppler spectra. Radio Sci. Vol.17, No. 3, PP. 643-652, 1982.
- [12] W. M. Huang, S. R. Wang, S. C. Wu, B. X. Wen, C. R. Qiu. Extraction of wind field from high-frequency radar sea-echo Doppler spectra. Journal of Wuhan University (Natural Science Edition), Vol. 45 No. 1, PP. 115-118, 1999.
- [13] R.K. Howell, J. Walsh. Measurment of ocean wave spectra using narrow-beam HF radar. IEEE J Oceanic Eng, Vol. 18 No. 3, PP. 296-305, 1993.
- [14] W. H. Shi, S. R. Wang, S. C. Wu,et al. Extraction of wave height parameters from high fre-quency radar sea echo Doppler spectrum. Journal of Wuhan University(Natural Science Edition), Vol. 44 No. 3, PP. 381-384,1998.